



**HAL**  
open science

## Drivers of soil carbon stabilization in oil palm plantations

Johanna Rüegg, Juan Carlos Quezada, Mathieu Santonja, Jaboury Ghazoul, Yakov Kuzyakov, Alexandre Buttler, Thomas Guillaume

► **To cite this version:**

Johanna Rüegg, Juan Carlos Quezada, Mathieu Santonja, Jaboury Ghazoul, Yakov Kuzyakov, et al.. Drivers of soil carbon stabilization in oil palm plantations. Land Degradation and Development, 2019, 30, pp.1904-1915. 10.1002/ldr.3380 . hal-02190829

**HAL Id: hal-02190829**

**<https://amu.hal.science/hal-02190829>**

Submitted on 23 Jul 2019

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Copyright

1 **TITLE: Drivers of soil carbon stabilization in oil palm plantations**

2

3 **RUNNING TITLE:** Soil carbon stabilization in oil palm plantations

4

5 **AUTHORS:** Johanna Rüegg<sup>1,2</sup>, Juan Carlos Quezada<sup>1,3</sup>, Mathieu Santonja<sup>1,3,4</sup>, Jaboury  
6 Ghazoul<sup>2</sup>, Yakov Kuzyakov<sup>5,6,7</sup>, Alexandre Buttler<sup>1,3,8</sup>, Thomas Guillaume<sup>1,3</sup>

7

8 **ADDRESSES**

9 <sup>1</sup>École Polytechnique Fédérale de Lausanne EPFL, School of Architecture, Civil and  
10 Environmental Engineering ENAC, Laboratory of ecological systems ECOS, 1015  
11 Lausanne, Switzerland.

12 <sup>2</sup>Chair of Ecosystem Management, Institute of Terrestrial Ecosystems, Department of  
13 Environmental Systems Science, ETHZ, 8092 Zürich.

14 <sup>3</sup>Swiss Federal Institute for Forest, Snow and Landscape Research WSL, Site Lausanne, 1015  
15 Lausanne, Switzerland.

16 <sup>4</sup>Aix Marseille Univ, Avignon Université, CNRS, IRD, IMBE, Marseille, France.

17 <sup>5</sup>Georg August University, Department of Soil Science of Temperate Ecosystems,  
18 Agricultural Soil Science, 37077 Göttingen, Germany.

19 <sup>6</sup>Institute of Environmental Sciences, Kazan Federal University, 420049 Kazan, Russia.

20 <sup>7</sup>Institute of Physicochemical and Biological Problems in Soil Science, Russian Academy of  
21 Sciences, 142290 Pushchino, Russia.

22 <sup>8</sup>Laboratoire de Chrono-Environnement, UMR CNRS 6249, UFR des Sciences et Techniques,  
23 16 route de Gray, Université de Franche-Comté, 25030 Besançon, France.

24

25 **CORRESPONDING AUTHOR:** Thomas Guillaume ([thomas.guillaume@epfl.ch](mailto:thomas.guillaume@epfl.ch))

26 **ABSTRACT**

27 Increasing soil organic carbon (SOC) in agroecosystems is necessary to mitigate climate  
28 change and soil degradation. Management practices designed to reach this goal call for a  
29 deeper understanding of the processes and drivers of soil carbon input stabilization. We  
30 identified main drivers of SOC stabilization in oil palm plantations using the well-defined  
31 spatial patterns of nutrients and litter application resulting from the usual management  
32 scheme. The stabilization of oil palm-derived SOC (OP-SOC) was quantified by  $\delta^{13}\text{C}$  from a  
33 shift of C4 (savanna) to C3 (oil palm) vegetations. Soil organic carbon stocks under frond  
34 piles were 20 and 22 % higher compared to harvest paths and interzones, respectively.  
35 Fertilization and frond stacking did not influence the decomposition of savanna-derived SOC.  
36 Depending on management zones, net OP-SOC stabilization equalled 16-27% of the fine root  
37 biomass accumulated for 9 years. This fraction was similar between frond piles and litter-free  
38 interzones, where mineral NPK fertilization is identical, indicating that carbon inputs from  
39 dead fronds did not stabilize in SOC. A path analysis confirmed that the OP-SOC distribution  
40 was largely explained by the distribution of oil palm fine roots, which itself depended on  
41 management practices. SOC mineralization was proportional to SOC content and was  
42 independent on phosphorus availability. We conclude that SOC stabilization was driven by C  
43 inputs from fine roots and was independent of alteration of SOC mineralization due to  
44 management. Practices favouring root growth of oil palms would increase carbon  
45 sequestration in soils without necessarily relying on the limited supply of organic residues.

46

47 **KEYWORDS:** carbon isotopes - fertilization - fine roots - microbial activity - structural  
48 equation modelling - savanna - Colombia

49

## 50 1. INTRODUCTION

51 Soil organic carbon (SOC) depletion in agroecosystems is a major source of greenhouse gas  
52 emissions, resulting in losses of soil fertility and ecosystem stability (Amundson et al., 2015).  
53 Policymakers have recently acknowledged the promotion of soil C sequestration in  
54 agroecosystems as a promising strategy to simultaneously mitigate climate change and  
55 enhance food security (Lal, 2016). Nonetheless, management practices favouring soil C  
56 sequestration lead to highly variable outcomes, calling for a deeper understanding of  
57 processes and factors controlling C stabilization, especially after land-use changes (Ghimire,  
58 Lamichhane, Acharya, Bista, & Sainju, 2017; Haddaway et al., 2017). This is particularly  
59 needed in regions that are undergoing rapid and substantial land degradation following  
60 conversion from largely forested landscapes to intensive agricultural systems, as has been the  
61 case in tropical landscapes dominated by oil palm (*Elaeis guineensis*) plantations. In Sumatra,  
62 a region with a long history of oil palm cultivation, a significant proportion of oil palm  
63 plantations has already reached a critical low level of SOC content (< 1%) in the topsoil  
64 (Guillaume, Holtkamp, Damris, Brümmer, & Kuzyakov, 2016).

65 Stocks of SOC depend on the balance between soil C inputs from vegetation and outputs  
66 from SOC mineralization, erosion and leaching (Lorenz & Lal, 2018). This balance is strongly  
67 affected by the conversion of natural ecosystems to intensive agricultural land (Guillaume et  
68 al., 2018). Predicting impacts of land-use change and management on SOC dynamics and its  
69 stabilization faces major difficulties as many factors affect both litter input and mineralization  
70 processes that determine the fraction of C input stabilized in SOC. While higher plant biomass  
71 inputs might lead to higher SOC accumulation, an increase of fresh organic matter inputs may  
72 enhance the mineralization of more recalcitrant SOC that ultimately reduces the gain in SOC,  
73 a process known as priming (Kuzyakov, Friedel, & Stahr, 2000). Nutrient application  
74 modifies the stoichiometry of organic matter inputs and of soil organic matter, which in turn

75 affects microbial processes controlling SOC stabilization (Qiao et al., 2016; Zang, Wang, &  
76 Kuzyakov, 2016). For instance, altered microbial carbon use efficiency (i.e. the ratio of C  
77 incorporated into microbial biomass to the added C) or mining for nutrients from recalcitrant  
78 SOC result in either SOC gains or losses (Finn et al., 2016; Kirkby et al., 2014). Quantifying  
79 soil C inputs remains a methodological challenge and few data are available, for instance, on  
80 root turnover, rhizodeposition, and the fraction of aboveground litter C stabilized in SOC  
81 (Pausch & Kuzyakov, 2018). Experimental data are especially limited for perennial plants in  
82 tropics and subtropics.

83         The fast SOC turnover in the tropics makes soils particularly sensitive to land-use  
84 change (Guillaume, Damris, & Kuzyakov, 2015; Pabst, Gerschlauer, Kiese, & Kuzyakov,  
85 2016; Zech et al., 1997). For example, soil C inputs decrease up to 90% when rainforests are  
86 converted to oil palm plantations, resulting in a rapid drop of SOC (Guillaume et al., 2018).  
87 Soil organic C losses are not, however, uniform within plantations, and specific management  
88 zones within the plantation may even exhibit a gain in SOC (Khasanah, van Noordwijk,  
89 Ningsih, & Rahayu, 2015; Rahman et al., 2018). Soil C inputs in mature oil palm plantations  
90 without cover crops (the most common practice) occurs mostly belowground through  
91 rhizodeposition because the understory is frequently cleared. Significant aboveground C  
92 inputs occur only under frond piles, i.e. zones where dead fronds are stacked. A gradient of  
93 fertilizer application is superimposed on the gradient of soil C input: most fertilizers are  
94 applied around the trunks and, in some cases, additionally to the whole surface area at lower  
95 rate, including the frond piles but excluding the harvest path. These management practises  
96 lead to characteristic management zones with specific factors affecting SOC dynamics (Fig  
97 1a). With plantation ageing, the heterogeneity of SOC distribution increases, depending on  
98 the distance to the tree, the presence of frond piles and fertilizer applications (Frazão,  
99 Paustian, Pellegrino Cerri, & Cerri, 2013; Goodrick et al., 2015). Carbon and nutrients cycling

100 are, thus, highly heterogeneous leading to a variable SOC equilibrium depending on  
101 management practices. This highlights the possibility to promote the increase of overall SOC  
102 stocks by redesigning management practices to reach this goal.

103 In Colombia, oil palm plantations are often established on native savanna grasslands  
104 dominated by grasses with a C4 photosynthesis pathway. The shift from C4 (grasses) to C3  
105 vegetation (oil palm) allows source determination in soil organic matter using its  $\delta^{13}\text{C}$   
106 signature (Balesdent & Mariotti, 1987). The aim of the present study is to disentangle the  
107 effects of soil C and nutrient inputs on newly accumulated SOC (oil palm-derived) and the  
108 decomposition of old SOC (savanna-derived), taking advantage of the specific patterns of  
109 fertilization and soil C inputs in the four management zones. We hypothesized that soil C  
110 inputs and consequent SOC accumulation increase with fine root density, which itself depends  
111 on plantation age and distance to trees, as well as on the presence of frond piles. Fertilizer  
112 application is, however, expected to decrease SOC stabilization rates where soil C inputs are  
113 low. Hence, the specific aims of the study are to i) quantify the new oil palm-derived (C3  
114 signature) and the old savanna-derived SOC stocks (C4) in a mature oil palm plantation  
115 established on native savanna grassland, ii) assess oil palm root development and its impact  
116 on SOC accumulation with plantation age, iii) determine the impacts of management practices  
117 on root development and soil microbial activity and iv) identify the main factors (root density,  
118 nutrient availability, microbial activity) driving SOC stabilization.

119

## 120 **2. MATERIALS AND METHODS**

121

### 122 ***2.1. Study area***

123 The study was conducted in the Eastern Plains (*Llanos Orientales*), Department Meta,  
124 Colombia (4°05'7.0"N, 71°53'59.0"W). The region experiences a tropical climate (mean

125 annual temperature of 26 °C and yearly precipitation of 2200 mm yr<sup>-1</sup>) with a distinct dry  
126 season from December to March and 95% of the yearly rain falling between April and  
127 November (Lavelle et al., 2014; Rippstein, Amézquita, Escobar, & Grollier, 2001). The study  
128 site lies in the slightly undulating well-drained high plains (Altillanura plana) dominated by  
129 Plinthosols and Ferralsols (IUSS Working Group WRB, 2014). These soils have a low  
130 fertility, high acidity and high aluminium saturation limiting agricultural production (Lavelle  
131 et al., 2014; Rippstein et al., 2001). The natural vegetation is an herbaceous savanna with  
132 scarce bushes, which is drained by many small rivers. Gallery forests (morichales) grow in  
133 the depressions along these rivers (Rippstein et al., 2001).

134 Two unmanaged native savanna sites and three oil palm plantations with increasing age  
135 (2-, 4-, and 9-year old) were selected within an area of approximately 8 × 8 km and sampled  
136 in July and August 2016 (Fig. S1). Sites were carefully chosen with the help of plantation  
137 agronomists to ensure that the investigated plantations had been established on unmanaged  
138 native savannas, i.e. no cattle grazing in the past. Soils in the five selected sites were classified  
139 as sandy-loam Ferralsols (clay content and pH ranging from 11 to 16% and from 4.5 to 5.3 in  
140 the top 30 cm, respectively) with compacted top soil (bulk density ranging from 1.31 to 1.46  
141 g cm<sup>-3</sup> in the top 10 cm). Drivers of SOC stabilization were assessed in the 9-year old  
142 plantation. The 2- and 4-year old plantations were selected to assess the development of oil  
143 palm rooting system and its relationship with SOC accumulation.

144

## 145 ***2.2. Plantation management***

146 Oil palm plantations had been established on native savannas whose soils were first loosened  
147 with a chisel plough (to a depth of 40 cm) and tilled with an overturning plough (10 cm).  
148 Liming and phosphate rocks were applied before planting palms. Oil palm trees had been  
149 planted in a triangular grid pattern with a distance of 9 m between trees, leading to 143 palms

150 ha<sup>-1</sup> (Fig. 1b). In the young plantations (2- and 4-year old), a mixture of Kudzu (*Pueraria*  
151 *phaseoloides*) and Desmodium (*Desmodium heterocarpon* subsp. *ovalifolium*) cover crops  
152 (C3 vegetation) have been implemented after planting. Cover crops were never used in the  
153 mature plantation (9 years), and in this case the soil was always kept bare after planting with  
154 oil palms. Accordingly, C3-derived SOC in the 9-year old plantation are attributed solely to  
155 oil palm-derived SOC.

156 Oil palm management leads to four well-defined management zones in productive  
157 plantations (about 4 years after establishment). In the study region, the weeded circle (WC)  
158 around the oil palm trunk is always kept free of vegetation (Fig. 1). At young ages (2- and 4-  
159 year old), all fertilizers are applied in WC. Associated with the beginning of harvest after 4  
160 years, pruning starts and fronds are piled up in between palm trees (frond pile, FP). Each  
161 second avenue between palm lines becomes a harvest path (HP), where machines circulate.  
162 In productive plantations (after 4 years), fertilizers are evenly spread from the harvest path by  
163 machines, i.e. all management zone receive the same amount of fertilizers except the harvest  
164 path that receive none. The remaining area, especially the alternating avenue, represents the  
165 fourth zone, where the soil is kept bare and fertilization starts after 4 years (interzone, IZ).  
166 The relative surface area of the four management zones were 60% (IZ), 18% (HP), 12% (WC)  
167 and 10% (FP). Fertilization depends on oil palm stand age. The 9-year old plantation received  
168 per hectare during the first year about 240 kg of NPK, 60 kg of kieserite (MgSO<sub>4</sub>), and 25 kg  
169 of KCl, of zinc and of boron. The amount of NPK increased over time while the use of other  
170 types of fertilizer varied from year to year. The year before sampling, 600 kg of NPK, 200 kg  
171 of phosphate rocks and 160 kg of KCl-MgO per hectare were spread in the plantation.

172

### 173 **2.3. Soil and roots sampling**

174 Plots of 1 ha were established in all plantations and savannas in areas with homogeneous soil,  
175 far from the influence of roads, rivers or groundwater, free of laterite formations (arecife) and  
176 without former amendment of compost or residues from processing oil mills. Five trees as  
177 replicates for management zones were selected randomly in each plantation. For each of the  
178 five replicate of trees, the adjacent management zones (FP, HP, IZ, and WC) were sampled  
179 on a systematic grid (Fig. 1). Frond piles, harvest paths and interzones were sampled at the  
180 same distance from trees (4.5 m) to assess the effects of management starting 4 years after  
181 establishment. Two additional points in the interzone were sampled at 2 and 3 m away from  
182 the trunk to assess the horizontal expansion of oil palm roots (IZ2, IZ3, Fig. 1). In the 2- and  
183 4-year old plantations, only the weeded circles and interzones were sampled because of the  
184 absence of frond piles and harvest path in young plantations. In the 2-year old plantation, the  
185 point IZ3 was not sampled, as root densities were already low in IZ2 (Fig. S2). In the savanna  
186 plots, five sampling points were selected at regular distances along a 100 m transect.

187         Roots and soils were sampled with a cylindrical corer of 5 cm diameter at three depth  
188 intervals (0-10, 10-20 and 20-30 cm). Roots were separated from soil by sieving at 2 mm and  
189 rinsed to remove attached mineral particles. Fine roots that passed through the sieve were  
190 manually picked. For the 2- and 4-year old plantations outside of weeded circles, cover crop  
191 roots were removed from oil palm roots. Oil palm roots were divided into coarse roots (> 2  
192 mm, corresponding to primary and secondary roots of oil palms) and fine roots (< 2 mm,  
193 corresponding to tertiary and quaternary roots of oil palms). Dry root biomass was determined  
194 after drying at 60 °C for 48 h. Soil samples were air-dried and sieved at 2 mm directly after  
195 collection and further oven dried at 40 °C for 48 h prior to laboratory analyses.

196

#### 197 **2.4. Soil analysis**

198 Total C and nitrogen (N) contents in soil, as well as  $\delta^{13}\text{C}$  signature were determined at the  
199 University of Göttingen with an isotope ratio mass spectrometer (Delta Plus, Finnigan MAT,  
200 Bremen, Germany). Because of the absence of carbonates in acidic soils, total C represents  
201 organic C. Residual water content was assessed by drying soil samples at 105 °C for 24 h.  
202 Bulk density was measured by inserting horizontally two cylinders of 100 cm<sup>3</sup> per depth at 0-  
203 5, 5-10, 10-20 and 20-30 cm depth in a soil pit located at the centre of each sampling plot.  
204 Bulk density in each cylinder was determined after drying at 105 °C and averaged between  
205 four cylinders for 0-10 cm depth interval and 2 cylinders for 10-20 and 20-30 cm depth  
206 intervals. Carbon stocks were calculated multiplying C contents with bulk density and the  
207 layer thickness and the respective fractions of C4 (savanna-derived C) and C3 (oil palm-  
208 derived) SOC. Soil available phosphorous (P) was determined using Bray II extraction  
209 method (Bray & Kurtz, 1945). Three grams of soil were extracted by shaking for 15 min with  
210 20 ml of 0.03 N NH<sub>4</sub>F and 0.025 N HCl. Filtered extracts were mixed with a colorimetric  
211 reagent ((NH<sub>4</sub>)<sub>6</sub>Mo<sub>7</sub>O<sub>24</sub> – SnCl<sub>2</sub>) and absorbance was measured with a UV/VIS spectrometer  
212 at 660 nm (Lambda 35, Perkin Elmer, Buckinghamshire, United Kingdom).

213

## 214 **2.5. C3 and C4 derived carbon**

215 Relative portion of oil palm- and savanna-derived C in SOC were calculated based on the  
216 differences in  $\delta^{13}\text{C}$  signature of biomass between savanna grassland, dominated by C4  
217 photosynthetic pathway, and oil palms, C3 photosynthetic pathway, using two sources linear  
218 isotopic mixing model (Balesdent & Mariotti, 1987):

$$219 \quad f_{OP\_d} = \frac{\delta^{13}\text{C}_{S\_d} - \delta^{13}\text{C}_{NS\_d}}{\delta^{13}\text{C}_{OP\_b} - \delta^{13}\text{C}_{NS\_b}}$$

220 where  $f_{OP\_d}$  is the fraction of oil palm (C3)-derived SOC,  $\delta^{13}\text{C}_{S\_d}$  is the isotopic signature  
221 measured in a soil sample at depth d,  $\delta^{13}\text{C}_{NS\_d}$  is the average isotopic signature at the

222 corresponding depth in natural savanna reference sites,  $\delta^{13}\text{C}_{\text{OP}_b}$  is the averaged signature of  
223 oil palm fine roots in the mature oil palm plantation (mean = -28.2‰, standard deviation (SD)  
224 = 0.18, n = 9) and  $\delta^{13}\text{C}_{\text{NS}_b}$  is the average isotopic signature of the aboveground and  
225 belowground savanna biomass determined on representative subsamples of aboveground  
226 biomass collected on 1 m<sup>2</sup> and fine roots collected with soil cores at each sampling point  
227 (mean = -13.6 ‰, SD = 1.1, n = 19). This approach assumes that the <sup>13</sup>C fractionation  
228 occurring during the integration of biomass into SOC at each soil depth is the same for  
229 savanna and oil palm biomass (Pausch & Kuzyakov, 2012). No <sup>13</sup>C fractionation was  
230 observed between savanna plant biomass and savanna SOC in 0-10 cm layer (-13.7 ‰, SD =  
231 0.5, n = 10).

232

### 233 ***2.6. Carbon stabilization per cumulative standing root biomass***

234 Fine root stocks served as a proxy for the C input from oil palm roots. To account for  
235 differences in root stocks between management zones and root development time depending  
236 on the distance to the palm tree, cumulative standing fine root biomass in each management  
237 zone was estimated by fitting a linear model on root biomass measured in various plantation  
238 ages. The model was integrated starting from plantation establishment for WC, but only  
239 starting from 4 years for IZ, FP and HP, i.e. when oil palm rooting system reached 4.5 m away  
240 from the tree (Fig. S2). Assuming constant belowground C input per unit of fine roots (Pausch,  
241 Tian, Riederer, & Kuzyakov, 2013), oil palm-derived SOC stocks were normalized per unit  
242 of cumulated fine root biomass to assess the effect of management zones on net C3  
243 stabilization efficiency.

244

### 245 ***2.7. Soil incubation***

246 Microbial biomass was analysed after incubation of topsoils (0-10 cm) from the 9-year old  
247 plantation (zones FP, HP, IZ and WC). Twenty grams of dry soil were rewetted to 60% WHC  
248 and incubated at 25 °C for 31 days. Jars were ventilated and weighed every week and rewetted  
249 once to compensate for the evaporated water. Carbon and nitrogen (N) in microbial biomass  
250 were measured at the end of the incubation by the fumigation-extraction method (Vance et  
251 al., 1987). For C and N contents, 5 g of incubated soil were fumigated for 24 h with ethanol-  
252 free CHCl<sub>3</sub> in a desiccator. Soils were extracted by shaking 1 h in 25 ml solution of 0.5 M  
253 K<sub>2</sub>SO<sub>4</sub> and then filtered. Non-fumigated samples were processed in parallel. Total extractable  
254 organic C and N were analysed with a TOC-N analyser (Shimadzu, Kyoto, Japan). Extractable  
255 C in the non-fumigated samples was assumed to represent dissolved organic carbon (DOC).  
256 Microbial biomass C and N were calculated as the difference between fumigated and non-  
257 fumigated samples, which were used also to calculate microbial C:N ratio. Microbial biomass  
258 C was corrected by dividing extractable C with a factor of 0.45 (Beck et al., 1997), N with a  
259 factor of 0.54 (Brookes, Landman, Pruden, & Jenkinson, 1985).

260

## 261 **2.8. Basal respiration**

262 Basal respiration of the rewetted samples was measured with the MicroResp™ kit (Campbell,  
263 Chapman, Cameron, Davidson, & Potts, 2003). Three analytical replicates of 0.5 g for each  
264 field replicate were taken from the incubation jars after rewetting and incubated in 96 deep-  
265 well plate in parallel to the incubation in jars. Soil was kept moist by a moist paper towel  
266 fixed on the plate. Respiration was measured 1, 2, 4, 8, 15, 18, 24 and 31 days after rewetting  
267 using a MicroResp™ kit (Campbell et al., 2003). To remove residual CO<sub>2</sub> in the wells, the  
268 plate was aerated with a fan before incubating for 6 h with the indicator plate on top (at 27  
269 °C). The indicator plate was read before and after incubation with a spectrophotometer  
270 (Microplate reader BioTek SynergyMX) at 570 nm. Absorption calibration was done by

271 dissolving a known amount of NaCO<sub>3</sub> with 1 M HCl in excess in closed jars with eight  
272 microwells of the indicator plate for 6 h (Campbell et al., 2003).

273 Soil respiration was partitioned using a two-pool mixed-model to describe SOC  
274 mineralization kinetics (Bonde & Lindberg, 1988). The first pool follows a first-order  
275 decomposition kinetics, while the second follows a zero-order kinetics, corresponding to the  
276 stabilized basal respiration:

$$277 \quad C_{min} = C_l(1 - e^{-tk_l}) + BR t$$

278 where  $t$  is the time,  $C_{min}$  is the cumulative CO<sub>2</sub> mineralized to time  $t$ ,  $C_l$  is the labile C pool  
279 released from sample preparation and re-wetting,  $k_l$  is the decomposition constant of the labile  
280 pool and  $BR$  is the basal respiration. The basal respiration of each field replicate was  
281 determined by fitting the model on the three analytical replicates. The metabolic quotient is  
282 the ratio of basal respiration over microbial biomass C ( $C_{mic}$ ).

283

## 284 **2.9. Statistical analyses**

285 All statistical analyses were performed using the open source software R version 3.2.1 (R  
286 Core Team, 2016). One sampling point in WC and one in IZ were removed from all analyses  
287 due to very low  $\delta^{13}C$  signatures, probably resulting from the former presence of C3 bush at  
288 these exact sampling locations. We used a linear mixed-effects model approach (“lme4”  
289 package), followed by Tukey HSD tests for post hoc pairwise comparisons, to test for the  
290 effects of management zones (FP, HP, IZ and WC) on soil parameters (bulk, C3 and C4  
291 stocks, fine root biomass, net C3 stabilization, basal respiration, microbial biomass C and C:N  
292 ratio, and metabolic quotient) at each soil depth (0-10, 10-20 and 20-30 cm) in the mature  
293 plantation (9-year old) with palm trees as random factor. Normal distribution of residuals and  
294 homogeneity of variance were tested by Shapiro and Levene tests, respectively, and data was  
295 log-transformed if necessary. Causal relationships between parameters (C stocks vs. age, C3

296 stocks vs. fine roots) were assessed by linear regressions. Average C stocks at plot scale down  
297 to 30 cm depth were calculated using the relative surface area of each management zone.  
298 Associations among parameters were calculated using Pearson correlation. *P-values* were  
299 determined using the function *cor.test*.

300 A path analysis was performed to disentangle the direct and indirect effects of fine roots,  
301 soil microorganisms and nutrient application on the accumulation of oil palm-derived SOC  
302 by using the “lavaan” package (Rosseel, 2012). Based on the priori knowledge, we developed  
303 an initial conceptual model that was both consistent with our data and which made biological  
304 sense (Fig. S3). We first created five conceptual groups of measured variables, which  
305 represented i) nutrient application (measure of available P), ii) fine roots C inputs (measure  
306 of fine roots biomass), iii) SOM quality (measure of C:N ratio of SOM), iv) SOC  
307 accumulation (measure of oil palm-derived SOC) and v) soil microorganism effects. As a  
308 proxy for the microbial effect, we used the score of each sample on the first axis (PC1 = 68  
309 %) of a principal component analysis (PCA) including all microbial related variables (soil  
310 basal respiration, metabolic quotient,  $N_{mic}$ ,  $C_{mic}$ , microbial C:N ratio) (Fig. S4). The  
311 conceptual model hypothesized that fine roots biomass and microbial activity have a direct  
312 impact on oil-palm SOC accumulation. Both might be affected by nutrient application.  
313 Additionally, microbial activity would be directly affected by root density through the amount  
314 of rhizodeposition and indirectly by changing organic matter quality. Finally, P availability  
315 could have a direct effect on microorganisms and SOM quality by changing the resources’  
316 stoichiometry in the soil. Pedoclimatic factors affecting SOC stabilization (e.g. soil  
317 mineralogy) were not specified in the model because they are identical between zones and  
318 cannot be modified by management. The adequacy of the model was determined by non-  
319 significant differences between the predicted and observed covariance matrices ( $\chi^2$  tests,  $p >$   
320 0.05), low root mean squared error of approximation index (RMSEA  $<$  0.1), high Tucker-

321 Lewis index (TLI > 0.90) and high comparative fit index (CFI > 0.90) (Grace, 2006; Rosseel,  
322 2012).

323

### 324 **3. RESULTS**

325

#### 326 ***3.1. Soil organic carbon stocks and origin***

327 Observations done on the 9-year old plantation showed lower soil C stocks compared to the  
328 native savanna grasslands (Fig. 2). On average,  $1.0 \pm 0.2$  kg C m<sup>-2</sup> was lost down to 30 cm  
329 depth, considering the relative area of each management zone. However, spatial SOC  
330 distribution in the mature oil palm plantation depended on management zones. Soil C losses  
331 during that period of time in the top 10 cm under frond piles and, to a lesser extent, under  
332 weeded circles were lower than C losses in harvest paths and interzones (Fig. 2). This trend  
333 was similar down to 30 cm depth but the differences were not significant below 10 cm depth.  
334 Differences in SOC stocks between management zones arose mainly from a higher  
335 accumulation of oil palm-derived SOC (C3-derived) under frond piles and weeded circles  
336 (Fig. 2). Oil palm-derived SOC after 9 years already accounted for between 27% (IZ) and  
337 45% (FP and WC) of the total SOC stock in the top 10 cm, where differences between zones  
338 were highest. Below 20 cm depth, the contribution of oil palm-derived SOC dropped,  
339 accounting for 2% (IZ) and maximum 12% (WC) of the total SOC stocks. The amount of  
340 savanna-derived SOC (C4-derived) remaining after 9 years was lower in the top 10 cm than  
341 between 10-20 cm depth, indicating a faster decomposition of this C pool in the top soil  
342 compared to deeper soil layers. Nonetheless, management zones had little influence on the  
343 decomposition rates of savanna-derived SOC, except in 10-20 cm depth under weeded circles  
344 where more C was lost than under frond piles. A similar trend was observed in the top 10 cm  
345 that explains why total SOC under weeded circles was intermediate as compared to frond

346 piles and harvest paths, despite the high amount of oil palm-derived SOC stabilized under  
347 weeded circles.

348

### 349 **3.2. Root development and C inputs**

350 Oil palm fine roots were first observed at 4.5 m away from palm trees in two out of the five  
351 investigated palms in the 4-year old plantations (Fig. S2). This indicates that oil palm rooting  
352 systems from adjacent palms started to overlap at that age, but root biomass at 4.5 m was still  
353 very low. Fine root biomass under weeded circles increased constantly during 9 years (Fig.  
354 3a). At the age of 9 years, roots were observed in all management zones but root growth was  
355 strongly enhanced under frond piles, reaching the same fine root biomass as under weeded  
356 circles in only 5 years, well above fine root biomass in the interzones and harvest paths (Fig.  
357 3a).

358 Oil palm-derived SOC under weeded circles was highly correlated with fine root  
359 biomass (Fig. 3b). Oil palm-derived SOC stocks corresponded to  $70 \pm 5$  % of fine root  
360 biomass stocks (slope = 0.70,  $R^2 = 0.80$ ) at the time of measurement, independently of soil  
361 depth and plantation age. This percentage was relatively constant between the three soil  
362 depths; from  $62 \pm 10$ % in the top 10 cm to  $50 \pm 11$ % between 20-30 cm depth (Table 1). Oil  
363 palm-derived SOC accumulation per year was 3 to 4 times faster in the top soil 10 cm than in  
364 the underlying layers (Table 1).

365 While oil palm-derived SOC in the weeded circle was accumulating from the beginning  
366 of the plantation, root biomass was measured at fixed time points. Oil palm fine root biomass  
367 was integrated over the whole duration of the plantation to calculate the cumulative fine roots  
368 biomass for each year and soil depth under weeded circles. Accordingly, oil palm-derived  
369 SOC accumulation corresponded to  $14 \pm 2$ %,  $11 \pm 2$ % and  $11 \pm 1$ % of the cumulated fine roots  
370 biomass stocks under weeded circles at 0-10, 10-20 and 20-30 cm depth, respectively (Table

371 1). Since soil depth had little impact on the relationship between oil palm-derived SOC and  
372 fine root biomass, differences in oil palm-derived SOC stabilization rates between depths  
373 were mainly related to differences in fine root biomass. The linear relationship across age  
374 showed that the proportion of oil-palm derived SOC stabilizing per amount of fine roots  
375 present in the plantation remains constant at least during 9 years (Fig. 3b).

376 Oil palm derived-SOC and fine root biomass were similar under frond piles and weeded  
377 circles in the top 10 cm in the 9-year old plantation (Fig. 3a). However, fine roots appeared  
378 at 4.5 m away from palms 4 years later than under weeded circles. To remove the effects of  
379 root density and duration of C inputs between zones, oil palm-derived SOC stocks were also  
380 divided by the cumulative roots biomass stocks and compared between management zones  
381 (Fig. 4).

382 After this normalization to the amount of fine root biomass, the stabilization of oil palm-  
383 derived SOC was similar between frond piles and interzones, two zones receiving the same  
384 amount of mineral fertilizers and sampled at the same distance to the tree. Consequently, only  
385 little C from the large amount of C present in dead fronds is eventually stabilized in SOC. The  
386 stabilization of oil palm-derived SOC was lower under weeded circles, the zone receiving the  
387 largest amount of mineral fertilizer, than under harvest paths, the only zone experiencing no  
388 direct application of mineral fertilizers.

389

### 390 ***3.3. Soil microorganisms and fertility***

391 Microbial activity (basal respiration and metabolic quotient) and biomass (microbial biomass  
392 C and microbial C:N ratio) parameters were similar under frond piles and weeded circles (Fig.  
393 5). In these two zones, the same amount of microbial biomass respired more C compared to  
394 harvest paths and interzones. Consequently, the metabolic quotient was highest under frond  
395 piles and weeded circles.

396 Available P (Bray II) was an order of magnitude higher in weeded circles compared to  
397 the other zones, as expected by the higher fertilization application in that zone. Weeded circles  
398 also had the lowest amount of K<sub>2</sub>SO<sub>4</sub>-extractable C (DOC) despite high SOC content. (Table  
399 S1). Between the three zones (HP, IZ and FP) located at the same distance to the palm trees  
400 but varying in their management, P availability was two times lower in harvest paths ( $11.1 \pm$   
401  $0.3 \mu\text{g g}^{-1}$ ) as compared to frond piles, but well above the P availability in native savanna sites  
402 ( $2.2 \pm 0.3 \mu\text{g g}^{-1}$ ), despite the absence of direct fertilization. The DOC amount was similar  
403 between the three zones and the C:N ratio was 1 unit higher under frond piles, indicating that  
404 only small change in SOC quality occurred between management zones.

405

#### 406 ***3.4. Drivers of the accumulation of oil palm-derived SOC***

407 In the three zones located at the same distance to the palm trees (HP, IZ and FP), oil palm-  
408 derived SOC stocks (C<sub>3</sub>) were highly correlated with fine root biomass ( $r = 0.82$ ), as well as  
409 to most microbial parameters, the amount of soil organic matter (C and N contents) and  
410 available P (Fig. 6). The basal respiration, the metabolic quotient and the microbial biomass  
411 N increased with higher oil palm-derived SOC but not with total C content, underlying the  
412 role of fresh organic C to maintain microbial activity. By contrast, higher C:N ratio in  
413 microbial biomass and, to a lesser extent, higher microbial biomass were associated to less  
414 oil palm-derived SOC and less fine roots biomass but with more savanna-derived SOC  
415 remaining after conversion to oil palm.

416 A path analysis was performed to disentangle the direct and indirect effects of fine roots,  
417 soil microorganisms and nutrient availability on the accumulation of oil palm-derived SOC  
418 at 4.5 m away from palm trees in the 9-year old plantation. Because of the high association  
419 among all microbial parameters (Fig. 6), the scores of each sample on the first axis (PC1 = 68  
420 %) of a principal component analysis (PCA) including all microbial related parameters were

421 used as a proxy for soil microorganisms in the path analysis (Fig. S4). The fitting parameters  
422 of the model were good ( $P\chi^2 = 0.33$ , RMSEA = 0.09, TLI = 0.97, CFI = 0.99), and the model  
423 explained 71% of the variance in oil palm-derived SOC (Fig. 7). Fine root biomass was  
424 strongly influenced by nutrient availability (available P) and was an important driver of  
425 microbial properties (i.e. increase in microbial biomass N, basal respiration, specific  
426 respiration and metabolic quotient and, in the opposite, decrease of microbial biomass C and  
427 C:N ratio – see also Fig. 6). By contrast, nutrient availability had only a marginal ( $p = 0.09$ )  
428 and opposite effect on soil microorganisms. While fine root biomass influenced SOM quality  
429 (C:N ratio), microorganisms were not affected by SOM quality. The direct effect of fine roots  
430 on oil palm-derived SOC stabilization was 2.8 times stronger than the marginal effect ( $p =$   
431  $0.06$ ) of soil microorganisms. Accordingly, nutrient availability had an important indirect  
432 impact on the accumulation of oil palm-derived SOC by favouring root development and  
433 thereby C inputs without enhancing SOC mineralization.

434

## 435 **4. DISCUSSION**

436

### 437 ***4.1. Drivers of SOC stabilization***

438 Soil organic C stocks strongly varied depending on the management zones of the plantation.  
439 Specific management impacted the amount of new oil palm C input and its stabilization into  
440 SOC but had little effect on the decomposition rate of old savanna-derived SOC (Fig. 2).  
441 Carbon stabilization was mainly driven by C inputs from fine roots rather than by changes in  
442 C outputs from microbial mineralization (Fig. 7). The relationship between fine roots and oil  
443 palm derived SOC remained weakly affected by soil depth, plantation age and management  
444 zones (Table 1, Fig. 3 and 4). Fine root biomass was a good proxy for soil belowground C  
445 inputs. Oil palm fine roots absorb nutrient and water, and therefore have definite growth and

446 short-term self-pruning. In contrast, coarse roots, which have the function of conduction, have  
447 indefinite growth and long term self-pruning (Jourdan, Michaux-Ferrière, & Perbal, 2000;  
448 Jourdan & Rey, 1997). The relationship between cumulated fine root biomass and oil-palm  
449 derived SOC stocks would not remain constant once stocks reach equilibrium. The  
450 relationship, however, was linear (Fig. 3), indicating that SOC stocks were still far from  
451 equilibrium after 9 years of cultivation and that oil palm-derived SOC will continue to  
452 increase.

453       Microbial metabolism and biomass were strongly affected by fine roots. Higher organic  
454 C availability in terms of oil palm-derived SOC and C inputs under frond piles and in the  
455 weeded circle were associated with microbial communities characterized by a high  
456 mineralization activity but of low efficiency – the so termed *r* strategy (Loeppmann,  
457 Blagodatskaya, Pausch, & Kuzyakov, 2016). The high metabolic quotient results from either  
458 a low C use efficiency or a high microbial biomass turnover. By contrast, zones with low C  
459 availability such as the harvest path and the interzones were associated with microbial  
460 communities more efficient to maintain their biomass despite lower mineralization rates.  
461 Their lower metabolic quotients and their higher C:N ratios suggest that the scarcity of C  
462 increased the proportion of K strategists and fungi within microbial communities (Mouginot  
463 et al., 2014; Six, Frey, Thiet, & Batten, 2006). While the main effect of nutrient availability  
464 on microbial communities was indirect by increasing roots C inputs, it tended to have also a  
465 minor direct but contrasting effect on microbial communities. Lower C to nutrients ratio  
466 favours high C use efficiency of microorganisms, which would explain the larger microbial  
467 biomass in interzones despite the low microbial activity and root density (Sinsabaugh,  
468 Manzoni, Moorhead, & Richter, 2013). Nonetheless, management impacts on microbial  
469 communities and their resource consumption strategies levelled-off resulting in similar  
470 fraction of SOC mineralized in all management zones. Weeded circles were the only area of

471 the plantation showing slightly lower net SOC stabilization and savanna-derived SOC stocks  
472 (Fig. 2 and 4). The very high amount of fertilizer applied already in the early stage of the  
473 plantation might have slightly fastened SOC turnover. Nonetheless, management effects were  
474 small in regards to the 10-fold increase of P availability in this zone as compared to the rest  
475 of the plantation. The fact that nutrient availability had little impact on the SOC mineralization  
476 and stabilization indicates that microorganisms were mostly C-limited. Indeed, soils were  
477 depleted in SOC and the whole surface area of the plantations, even areas not directly  
478 fertilized, experienced an increase in P availability as compared to native savannas.

479 In summary, management and its impact on soil belowground C inputs and nutrient  
480 availability did not lead to priming of recalcitrant SOC nor to faster turnover of fresh organic  
481 SOC, except around palms in the weeded circles. Hence, SOC distribution was not driven by  
482 an altered decomposition of SOC pools but by different rates of fine root growth and the  
483 resulting soil C inputs depending on management.

484

#### 485 ***4.2. Drivers of soil C inputs***

486 Oil palm rooting system follows a relatively rigid and genetically determined development  
487 (Jourdan et al., 2000; Jourdan & Rey, 1997). Primary lateral roots start to grow one year after  
488 germination at a rate of 3 mm d<sup>-1</sup>, confirming that roots of palms, which are planted a year  
489 after germination, reach 4.5 m in the fourth year of a plantation (Jourdan & Rey, 1997).  
490 Nevertheless, root development showed plasticity, reacting to management as shown by the  
491 fast development of fine roots under frond piles. Root development is generally enhanced in  
492 nutrients-rich zones (Hodge, 2004). Phosphorous availability was an important driver of root  
493 development, favoring soil C inputs from fine roots, and consequently the accumulation of  
494 new SOC (Fig. 7). Its distribution did not exactly reflect the pattern of fertilizer application  
495 (Table S1). Despite no direct fertilization applications, harvest paths were enriched in

496 available P as compared to the reference savanna sites and only frond piles exhibited a  
497 significantly higher P availability. This discrepancy may result from the initial soil  
498 preparation with dolomite and phosphate rocks and seems to be maintained over time by the  
499 recycling of organic P inputs from rhizodeposition, root turnover, and frond mineralization.  
500 The positive impact of frond piles on root development is likely not limited to P availability.  
501 Fronds piles increase the availability of major cations (Law, Husni, Ahmed, & Haniff Harun,  
502 2009). Similarly, it was shown that the application of empty fruit bunches on the surface  
503 increases soil moisture (Tao, Slade, Willis, Caliman, & Snaddon, 2016) and roots  
504 development (Kheong, Rahman, Musa, & Hussein, 2010). The positive impact of frond piles  
505 on SOC stocks has been previously reported (Haron, Brookes, Anderson, & Zakaria, 1998;  
506 Law et al., 2009) but authors have already highlighted the small increase of SOC stocks given  
507 the huge amount of frond's biomass C concentrated on a small surface area ( $2-3 \text{ kg C m}^{-2} \text{ yr}^{-1}$   
508 <sup>1</sup> in frond piles). It was suggested that fronds were mineralized mostly aboveground with little  
509 contribution to SOC (Haron et al., 1998). Our findings confirm this hypothesis and indicate  
510 that the positive impacts of frond piles arise more from the improvement of soil conditions  
511 (likely nutrient availability, humidity, protection from erosion) that favours root growth than  
512 from their role as a C source.

513

#### 514 ***4.3. Increasing SOC stocks***

515 Soil organic C stabilization depends on belowground C inputs and not on the management  
516 induced variation of SOC mineralization rates (Fig. 7). This field evidence-based finding is  
517 similar to the conclusion of a modelling study on arable cropping system in temperate zones  
518 (Autret et al., 2016). Soil C inputs should be enhanced to increase soil C sequestration in oil  
519 palm plantations. The current management practice that consist in piling dead fronds,  
520 however, is not efficient to integrate the fronds' organic matter into the soil. This can be

521 explained by several mechanisms. First, organic matter quality in fronds is low due to high C  
522 to nutrient ratios (Yusuyin et al., 2015). This decreases the C use efficiency of  
523 microorganisms. Second, bioturbation from soil fauna is limited as indicated by a sharp  
524 transition between decomposing fronds and the soil surface observed in the field. Soil fauna  
525 abundance, especially of earthworms, and biogenic macroaggregates are lower in oil palm  
526 plantations as compared to native savanna and in improved pastures (Lavelle et al., 2014).  
527 Consequently, fronds' organic matter does not benefit from the protection mechanisms that  
528 minerals would provide if it was integrated into the soils by fauna activity (Schmidt et al.,  
529 2011). Application of composted mill residues would be a solution to decrease C to nutrient  
530 ratios of litter and improve the integration of organic matter into the soil by favouring soil  
531 fauna activity, humification and mineral protection.

532         The availability of organic residues from the palm oil production chain that could be  
533 applied in plantations, however, is limited. Leguminous cover crops, as already implemented  
534 in the younger plantations, are an alternative to increase aboveground and belowground C and  
535 N inputs. Enhancing root growth by mimicking the effects of frond piles is a promising  
536 solution to increase SOC if palms have plasticity to allocate more C to their rooting system  
537 when soil conditions are favourable or if palm varieties are developed for that purpose. Oil  
538 palm fronds could be spread on larger surface area or mixed with other residues, such as  
539 empty fruit bunched or fibers, to make a mulch. Even solutions not based on organic matter  
540 that would limit soil evaporation, retain nutrients, limit run-offs and erosion might improve  
541 root development. Future research should address whether SOC stocks eventually recover to  
542 initial SOC levels. Ensuring long-term soil fertility in oil palm plantations is fundamental to  
543 avoid a conversion of natural ecosystem constrained by soil degradation in older plantation.  
544 The benefits from increasing SOC in terms of climate change mitigation would cascade far  
545 beyond the amount of C sequestered in the soil because of the gain in soil fertility.

546

547 **ACKNOWLEDGEMENTS**

548 This study was financed by the Swiss National Science Foundation (r4d – Ecosystems) “Oil  
549 Palm Adaptive Landscape” no. 152019. Contribution of YK was supported by the Russian  
550 Science Foundation (project No. 18-14-00362). We thank the owners and agronomists of the  
551 Palmeras Sillatava and Samani Italcol de Occidente for granting us access to their land and  
552 for their logistical support.

553

554 **CONFLICT OF INTEREST STATEMENT**

555 The authors declare no conflict of interest.

556

557 **REFERENCES**

- 558 Amundson, R., Berhe, a. a., Hopmans, J. W., Olson, C., Sztein, a. E., & Sparks, D. L. (2015).  
559 Soil and human security in the 21st century. *Science*, *348*, 1261071–1261071.  
560 <https://doi.org/10.1126/science.1261071>
- 561 Autret, B., Mary, B., Chenu, C., Balabane, M., Girardin, C., Bertrand, M., ... Beaudoin, N.  
562 (2016). Alternative arable cropping systems : A key to increase soil organic carbon  
563 storage? Results from a 16 year fi eld experiment. *Agriculture, Ecosystems and*  
564 *Environment*, *232*, 150–164. <https://doi.org/10.1016/j.agee.2016.07.008>
- 565 Balesdent, J., & Mariotti, A. (1987). Natural <sup>13</sup>C abundance as a tracer for studies of soil  
566 organic matter dynamics. *Soil Biology and Biochemistry*, *19*, 25–30.
- 567 Beck, T., Joergensen, R. G., Kandeler, E., Makeschin, F., Nuss, E., Oberholzer, H. R., &  
568 Scheu, S. (1997). An inter-laboratory comparison of ten different ways of measuring soil  
569 microbial biomass C. *Soil Biology and Biochemistry*, *29*, 1023–1032.  
570 [https://doi.org/10.1016/s0038-0717\(97\)00030-8](https://doi.org/10.1016/s0038-0717(97)00030-8)

- 571 Bonde, T. A., & Lindberg, T. (1988). Nitrogen Mineralization Kinetics in Soil During Long-  
572 term Aerobic Laboratory Incubations: A Case Study. *Journal of Environment Quality*,  
573 17, 414–417. <https://doi.org/10.2134/jeq1988.00472425001700030011x>
- 574 Bray, R. H., & Kurtz, L. T. (1945). Determination of total, organic, and available forms of  
575 phosphorous in soils. *Soil Science*, 59, 39–45. [https://doi.org/10.1097/00010694-](https://doi.org/10.1097/00010694-194501000-00006)  
576 194501000-00006
- 577 Brookes, P. C., Landman, A., Pruden, G., & Jenkinson, D. S. (1985). Chloroform fumigation  
578 and the release of soil-nitrogen - a rapid direct extraction method to measure microbial  
579 biomass nitrogen in soil. *Soil Biology and Biochemistry*, 17, 837–842.  
580 [https://doi.org/10.1016/0038-0717\(85\)90144-0](https://doi.org/10.1016/0038-0717(85)90144-0)
- 581 Campbell, C. D., Chapman, S. J., Cameron, C. M., Davidson, M. S., & Potts, J. M. (2003). A  
582 rapid microtiter plate method to measure carbon dioxide evolved from carbon substrate  
583 amendments so as to determine the physiological profiles of soil microbial communities  
584 by using whole soil. *Applied and Environmental Microbiology*, 69, 3593–3599.  
585 <https://doi.org/10.1128/aem.69.6.3593-3599.2003>
- 586 Finn, D., Page, K., Catton, K., Kienzle, M., Robertson, F., Armstrong, R., & Dalal, R. (2016).  
587 Ecological stoichiometry controls the transformation and retention of plant-derived  
588 organic matter to humus in response to nitrogen fertilisation. *Soil Biology and*  
589 *Biochemistry*, 99, 117–127. <https://doi.org/10.1016/j.soilbio.2016.05.006>
- 590 Frazão, L. A., Paustian, K., Pellegrino Cerri, C. E., & Cerri, C. C. (2013). Soil carbon stocks  
591 and changes after oil palm introduction in the Brazilian Amazon. *GCB Bioenergy*, 5,  
592 384–390. <https://doi.org/10.1111/j.1757-1707.2012.01196.x>
- 593 Ghimire, R., Lamichhane, S., Acharya, B. S., Bista, P., & Sainju, U. M. (2017). Tillage, crop  
594 residue, and nutrient management effects on soil organic carbon in rice-based cropping  
595 systems: A review. *Journal of Integrative Agriculture*, 16(1), 1–15.

596 [https://doi.org/10.1016/S2095-3119\(16\)61337-0](https://doi.org/10.1016/S2095-3119(16)61337-0)

597 Goodrick, I., Nelson, P., Nake, S., Webb, M., Bird, M., & Huth, N. (2015). Tree-scale spatial  
598 variability of soil carbon cycling in a mature oil palm plantation. *Soil Research*, *54*, 397–  
599 406. <https://doi.org/10.1071/SR15211>

600 Grace, J. B. (2006). *Structural equation modeling and natural systems*. Cambridge, UK:  
601 Cambridge University Press. <https://doi.org/10.1017/CBO9780511617799>

602 Guillaume, T., Damris, M., & Kuzyakov, Y. (2015). Losses of soil carbon by converting  
603 tropical forest to plantations: Erosion and decomposition estimated by  $\delta^{13}C$ . *Global*  
604 *Change Biology*, *21*, 3548–3560. <https://doi.org/10.1111/gcb.12907>

605 Guillaume, T., Holtkamp, A. M., Damris, M., Brümmer, B., & Kuzyakov, Y. (2016). Soil  
606 degradation in oil palm and rubber plantations under land resource scarcity. *Agriculture,*  
607 *Ecosystems & Environment*, *232*, 110–118. <https://doi.org/10.1016/j.agee.2016.07.002>

608 Guillaume, T., Kotowska, M. M., Hertel, D., Knohl, A., Krashevskaya, V., Murti Laksono, K.,  
609 ... Kuzyakov, Y. (2018). Carbon Costs and Benefits of Indonesian Rainforest  
610 Conversion to Plantations. *Nature Communications*, *9*, 2388.  
611 <https://doi.org/10.1038/s41467-018-04755-y>

612 Haddaway, N. R., Hedlund, K., Jackson, L. E., Kätterer, T., Lugato, E., Thomsen, I. K., ...  
613 Isberg, P. E. (2017). How does tillage intensity affect soil organic carbon? A systematic  
614 review. *Environmental Evidence*, *6*(1), 1–48. [https://doi.org/10.1186/s13750-017-0108-](https://doi.org/10.1186/s13750-017-0108-9)  
615 [9](https://doi.org/10.1186/s13750-017-0108-9)

616 Haron, K., Brookes, P. C., Anderson, J. M., & Zakaria, Z. Z. (1998). Microbial biomass and  
617 soil organic matter dynamics in oil palm (*Elaeis guineensis* jacq.) plantations, West  
618 Malaysia. *Soil Biology and Biochemistry*, *30*, 547–552. [https://doi.org/10.1016/S0038-](https://doi.org/10.1016/S0038-0717(97)00217-4)  
619 [0717\(97\)00217-4](https://doi.org/10.1016/S0038-0717(97)00217-4)

620 Hodge, A. (2004). The plastic plant: Root responses to heterogeneous supplies of nutrients.

621 *New Phytologist*, 162, 9–24. <https://doi.org/10.1111/j.1469-8137.2004.01015.x>

622 IUSS Working Group WRB. (2014). *World reference base for soil resources 2014. World*  
623 *Soil Resources Reports* (Vol. 106). FAO, Rome. Retrived from  
624 <http://www.fao.org/3/i3794en/I3794en.pdf>.

625 Jourdan, C., Michaux-Ferrière, N., & Perbal, G. (2000). Root system architecture and  
626 gravitropism in the oil palm. *Annals of Botany*, 85, 861–868.  
627 <https://doi.org/10.1006/anbo.2000.1148>

628 Jourdan, C., & Rey, H. (1997). Architecture and development of the oil-palm (*Elaeis*  
629 *guineensis* Jacq.) root system. *Plant and Soil*, 189, 33–48.  
630 <https://doi.org/10.1023/A:1004290024473>

631 Khasanah, N., van Noordwijk, M., Ningsih, H., & Rahayu, S. (2015). Carbon neutral? No  
632 change in mineral soil carbon stock under oil palm plantations derived from forest or  
633 non-forest in Indonesia. *Agriculture, Ecosystems & Environment*, 211, 195–206.  
634 <https://doi.org/10.1016/j.agee.2015.06.009>

635 Kheong, L. V., Rahman, Z. A., Musa, M. H., & Hussein, A. (2010). Empty fruit bunch  
636 application and oil palm root proliferation. *Journal of Oil Palm Research*, 22, 750–757.  
637 Retrived from [http://jopr.mpob.gov.my/wp-content/uploads/2013/09/jopr22april10-](http://jopr.mpob.gov.my/wp-content/uploads/2013/09/jopr22april10-Liew21.pdf)  
638 [Liew21.pdf](http://jopr.mpob.gov.my/wp-content/uploads/2013/09/jopr22april10-Liew21.pdf)

639 Kirkby, C. A., Richardson, A. E., Wade, L. J., Passioura, J. B., Batten, G. D., Blanchard, C.,  
640 & Kirkegaard, J. A. (2014). Nutrient availability limits carbon sequestration in arable  
641 soils. *Soil Biology and Biochemistry*, 68, 402–409.  
642 <https://doi.org/10.1016/j.soilbio.2013.09.032>

643 Kuzyakov, Y., Friedel, J. K., & Stahr, K. (2000). Review of mechanisms and quantification  
644 of priming effects. *Soil Biology and Biochemistry*, 32, 1485–1498.  
645 [https://doi.org/10.1016/s0038-0717\(00\)00084-5](https://doi.org/10.1016/s0038-0717(00)00084-5)

- 646 Lal, R. (2016). Beyond COP 21: Potential and challenges of the “4 per Thousand” initiative.  
647 *Journal of Soil and Water Conservation*, 71, 20A–25A.  
648 <https://doi.org/10.2489/jswc.71.1.20A>
- 649 Lavelle, P., Rodriguez, N., Arguello, O., Bernal, J., Botero, C., Chaparro, P., ... Fonte, S. J.  
650 (2014). Soil ecosystem services and land use in the rapidly changing orinoco river basin  
651 of colombia. *Agriculture, Ecosystems & Environment*, 185, 106–117.  
652 <https://doi.org/10.1016/j.agee.2013.12.020>
- 653 Law, M. C., Husni, M. H. A., Ahmed, O. H., & Haniff Harun, M. (2009). Spatial variability  
654 of soil organic carbon in oil palm: a comparison between young and mature stands.  
655 *International Journal of Agricultural Research*, 4, 402–417.
- 656 Loepmann, S., Blagodatskaya, E., Pausch, J., & Kuzyakov, Y. (2016). Substrate quality  
657 affects kinetics and catalytic efficiency of exo-enzymes in rhizosphere and detritusphere.  
658 *Soil Biology and Biochemistry*, 92, 111–118.  
659 <https://doi.org/10.1016/j.soilbio.2015.09.020>
- 660 Lorenz, K., & Lal, R. (2018). Soil Carbon Stock. In *Carbon Sequestration in Agricultural*  
661 *Ecosystems* (pp. 39–136). Springer International Publishing.  
662 [https://doi.org/10.1007/978-3-319-92318-5\\_2](https://doi.org/10.1007/978-3-319-92318-5_2)
- 663 Mougnot, C., Kawamura, R., Matulich, K. L., Berlemont, R., Allison, S. D., Amend, A. S.,  
664 & Martiny, A. C. (2014). Elemental stoichiometry of fungi and bacteria strains from  
665 grassland leaf litter. *Soil Biology and Biochemistry*, 76, 278–285.  
666 <https://doi.org/10.1016/j.soilbio.2014.05.011>
- 667 Pabst, H., Gerschlauser, F., Kiese, R., & Kuzyakov, Y. (2016). Land Use and Precipitation  
668 Affect Organic and Microbial Carbon Stocks and the Specific Metabolic Quotient in  
669 Soils of Eleven Ecosystems of Mt. Kilimanjaro, Tanzania. *Land Degradation &*  
670 *Development*, 27, 592–602. <https://doi.org/10.1002/ldr.2406>

- 671 Pausch, J., & Kuzyakov, Y. (2012). Soil organic carbon decomposition from recently added  
672 and older sources estimated by  $^{13}\text{C}$  values of  $\text{CO}_2$  and organic matter. *Soil Biology and*  
673 *Biochemistry*, 55, 40-47. <https://doi.org/10.1016/j.soilbio.2012.06.007>
- 674 Pausch, J., & Kuzyakov, Y. (2018). Carbon input by roots into the soil: Quantification of  
675 rhizodeposition from root to ecosystem scale. *Global Change Biology*, 24, 1–12.  
676 <https://doi.org/10.1111/gcb.13850>
- 677 Pausch, J., Tian, J., Riederer, M., & Kuzyakov, Y. (2013). Estimation of rhizodeposition at  
678 field scale: Upscaling of a  $^{14}\text{C}$  labeling study. *Plant and Soil*, 364, 273–285.  
679 <https://doi.org/10.1007/s11104-012-1363-8>
- 680 Qiao, N., Xu, X., Hu, Y., Blagodatskaya, E., Liu, Y., Schaefer, D., & Kuzyakov, Y. (2016).  
681 Carbon and nitrogen additions induce distinct priming effects along an organic-matter  
682 decay continuum. *Scientific Reports*, 6, 19865. <https://doi.org/10.1038/srep19865>
- 683 Rahman, N., De Neergaard, A., Magid, J., Van De Ven, G. W. J., Giller, K. E., & Bruun, T.  
684 B. (2018). Changes in soil organic carbon stocks after conversion from forest to oil palm  
685 plantations in Malaysian Borneo. *Environmental Research Letters*, 13, 105001.  
686 <https://doi.org/10.1088/1748-9326/aade0f>
- 687 Rippstein, G., Amézquita, E., Escobar, G., & Grollier, C. (2001). Condiciones naturales de la  
688 sabana. In *Agroecología y biodiversidad de las sabanas en los Llanos Orientales de*  
689 *Colombia* (pp. 1–21). Cali, Colombia: Centro Internacional de Agricultura Tropical.
- 690 Rosseel, Y. (2012). lavaan: An R package for structural equation modelling. *Journal of*  
691 *Statistical Software*, 48, 1–36. <https://doi.org/10.18637/jss.v048.i02>
- 692 Schmidt, M. W. I., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., Janssens, I. a., ...  
693 Trumbore, S. E. (2011). Persistence of soil organic matter as an ecosystem property.  
694 *Nature*, 478, 49–56. <https://doi.org/10.1038/nature10386>
- 695 Sinsabaugh, R. L., Manzoni, S., Moorhead, D. L., & Richter, A. (2013). Carbon use efficiency

696 of microbial communities: Stoichiometry, methodology and modelling. *Ecology Letters*,  
697 16, 930–939. <https://doi.org/10.1111/ele.12113>

698 Six, J., Frey, S. D., Thiet, R. K., & Batten, K. M. (2006). Bacterial and Fungal Contributions  
699 to Carbon Sequestration in Agroecosystems. *Soil Science Society of America Journal*,  
700 70, 555. <https://doi.org/10.2136/sssaj2004.0347>

701 Tao, H.-H., Slade, E. M., Willis, K. J., Caliman, J.-P., & Snaddon, J. L. (2016). Effects of soil  
702 management practices on soil fauna feeding activity in an Indonesian oil palm plantation.  
703 *Agriculture, Ecosystems & Environment*, 218, 133–140.  
704 <https://doi.org/10.1016/j.agee.2015.11.012>

705 Team, R. C. (2016). *R: A language and environment for statistical computing*. Vienna,  
706 Austria: R Foundation for Statistical Computing. Retrieved from [https://www.r-](https://www.r-project.org/)  
707 [project.org/](https://www.r-project.org/)

708 Vance , E. D., Brookes, P. C. & Jenkinson, D. S. (1987). An extraction method for measuring  
709 soil microbial biomass C. *Soil biology and Biochemistry*, 19(6), 703-707.  
710 [http://doi.org/10.1016/0038-0717\(87\)90052-6](http://doi.org/10.1016/0038-0717(87)90052-6)

711 Yusuyin, Y., Tan, N. P., Wong, M. K., Abdu, A. Bin, Iwasaki, K., & Tanaka, S. (2015).  
712 Nutrient Status of Frond Heaps and the Underlying Soils at An 18-Year-Old Oil Palm  
713 Field in Central Pahang, Malaysia. *Tropical Agriculture and Development*, 59, 212–220.  
714 <https://doi.org/10.11248/jsta.59.212>

715 Zang, H., Wang, J., & Kuzyakov, Y. (2016). N fertilization decreases soil organic matter  
716 decomposition in the rhizosphere. *Applied Soil Ecology*, 108, 47–53.  
717 <https://doi.org/10.1016/j.apsoil.2016.07.021>

718 Zech, W., Senesi, N., Guggenberger, G., Kaiser, K., Lehmann, J., Miano, T. M., ... Schroth,  
719 G. (1997). Factors controlling humification and mineralization of soil organic matter in  
720 the tropics. *Geoderma*, 79, 117–161. [https://doi.org/10.1016/S0016-7061\(97\)00040-2](https://doi.org/10.1016/S0016-7061(97)00040-2)

721 **TABLES**

722

723 **Table 1.** Oil palm-derived SOC stabilization under weeded circles at each depth depending  
 724 on time, fine root biomass in the plantation and fine root biomass cumulated since plantation  
 725 establishment (mean  $\pm$  SE). All linear regressions were significant at  $p < 0.001$ .

726

Weeded circles Depths <i>cm</i>	OP-derived SOC per year		OP-derived SOC per fine roots		OP-derived SOC per cumulative fine roots	
	<i>g SOC m<sup>-2</sup> yr<sup>-1</sup></i>	<i>R<sup>2</sup></i>	<i>g SOC g<sup>-1</sup> roots</i>	<i>R<sup>2</sup></i>	<i>g SOC g<sup>-1</sup> roots</i>	<i>R<sup>2</sup></i>
	0-10 (n=14)	54 $\pm$ 6	0.85	0.62 $\pm$ 0.10	0.76	0.14 $\pm$ 0.02
10-20 (n=15)	18 $\pm$ 6	0.55	0.53 $\pm$ 0.11	0.60	0.11 $\pm$ 0.02	0.59
20-30 (n=15)	12 $\pm$ 2	0.75	0.50 $\pm$ 0.11	0.57	0.11 $\pm$ 0.01	0.80

727

728

729 **FIGURE LEGENDS**

730

731 **Fig. 1.** Management zones in a mature oil palm plantation. **(a)** Four management zones  
732 varying in terms of fertilization and aboveground C inputs. **(b)** Sampling points (red dots) in  
733 the 9-year old plantation. Frond piles and harvest paths are absent in the 2- and 4-year old  
734 plantations but two additional points (empty red dots) were sampled to assess palm root lateral  
735 extension

736

737 **Fig. 2.** Soil organic C stocks separated between oil palm-derived soil organic carbon (C3-  
738 derived SOC) and savanna-derived SOC (C4-derived C) depending on management zones  
739 and soil depths after 9 years of oil palm cultivation on savanna grasslands. Mean values  $\pm$  SE  
740 are represented (n = 5 for FP and HP, n = 4 for IZ and WC). FP, frond pile; HP, harvest path;  
741 IZ, interzone; WC, weeded circle. Letters indicate significant differences between  
742 management zones in total SOC (upper-case) and oil palm-derived SOC (black lower-case)  
743 and savanna-derived SOC (white lower-case). The red continuous line shows the original  
744 SOC level ( $\pm$ SE) in savanna grasslands (n = 10). The difference between red line and the top  
745 of the stacked bars show the C losses over 9 years of oil palm cultivation

746

747 **Fig. 3.** Oil palm fine root development and oil palm-derived soil organic carbon (SOC) stocks.  
748 **(a)** Oil palm fine root development with plantation age in the top 10 cm under frond piles  
749 (FP), weeded circles (WC), harvest paths (HP), and interzones (IZ). Mean values  $\pm$  SE are  
750 represented (n = 5 for FP and HP, n = 4 for IZ and WC). Letters indicate significant differences  
751 between management zones in the 9-year old plantation; **(b)** relationship between oil palm-  
752 derived SOC (C3-SOC) and fine root biomass under weeded circles. Overall linear regression  
753 is indicated by the dashed line. Negative C3-SOC values result from the natural variation

754 standard of reference sites around the mean  $\delta^{13}C$  value when C3-SOC accumulation is very  
755 low or absent. Negative data were not set to zero to avoid increasing artificially the mean of  
756 the respective depth and age and thus decreasing the overall slope of the relationship. The two  
757 large arrows show the opposite effects of the plantation time and soil depth on the amount of  
758 new C (C3) stabilized in soil

759

760 **Fig. 4.** Net oil-palm derived SOC (C3) stabilized per amount of cumulated fine root biomass  
761 in each management zone after 9 years. Different letters indicate significant differences  
762 between harvest paths (HP), interzones (IZ), frond piles (FP), and weeded circles (WC). Mean  
763 values  $\pm$  SE are represented (n = 5 for FP and HP, n = 4 for IZ and WC)

764

765 **Fig. 5.** Relative effects of management zones on soil microorganisms in the 9-year old  
766 plantation: basal respiration, metabolic quotient, microbial biomass C, and C:N ratio of  
767 microbial biomass. Microbial variables in interzones (IZ), frond piles (FP), and weeded  
768 circles (WC) were normalized with their respective mean value in the harvest paths (HP), that  
769 is, the zone receiving neither C inputs nor nutrient applications. Specific respiration (basal  
770 respiration divided by C content) is not represented because the effect of management zone  
771 was not significant. Error bars represent SE (n = 5 for FP and HP, n = 4 for IZ and WC).  
772 Letters indicate significant differences between management zones for each parameter

773

774 **Fig. 6.** Pearson correlation matrix among soil variables in the top 10 cm. Correlation  
775 performed on samples collected in the three management zones located at 4.5 m away from  
776 palm trees (harvest paths, interzones, and frond piles). Variables are: basal respiration,  
777 specific respiration, metabolic quotient, oil palm-derived (C3), and savanna-derived (C4)  
778 SOC stocks, soil C and N contents (C, N), microbial biomass C and N (Cmic and Nmic), C:N

779 ratios of SOM (CN), K<sub>2</sub>SO<sub>4</sub>-extractable C (DOC), soil available P (P), and fine root biomass  
780 stocks (Fine roots). Only significant correlations are represented ( $p < .05$ ,  $n = 14$ )

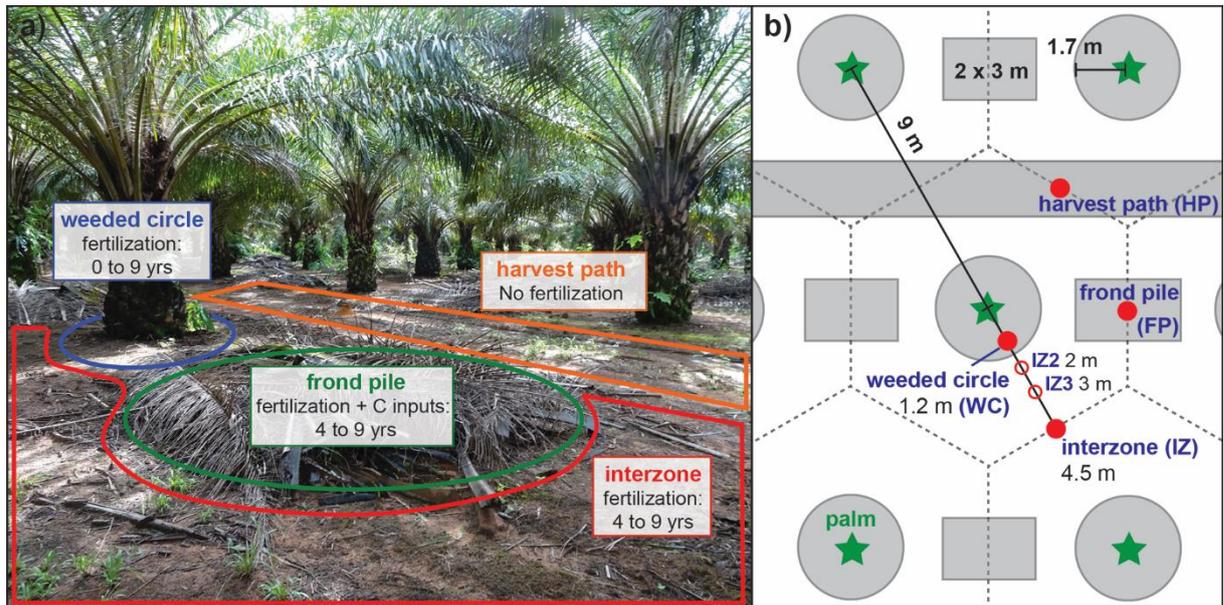
781

782 **Fig. 7.** Drivers of oil palm-derived soil organic carbon (SOC) accumulation in the top 10 cm  
783 of the three management zones located at 4.5 m away from palm trees (harvest paths,  
784 interzones, and frond piles). Phosphorus availability was used as proxy for nutrient  
785 availability. Fine roots correspond to fine root biomass stocks. Scores of samples on the first  
786 axis (68%) of a principal component analysis (PCA) of all microbially related variables were  
787 used as proxy for microbial effects. The C:N ratio of SOM was used as proxy for soil organic  
788 matter quality. Solid arrows represent significant effects ( $*p < .05$ ,  $***p < .001$ ,  $n = 14$ ) and  
789 widths are proportional to the effect. Dashed arrows represent marginally effects ( $p < .10$ ).  
790 Non-significant relationships ( $p > .10$ ) are not represented (nutrient availability to SOM  
791 quality and SOM quality to soil microorganisms)

792

793 **Fig. 1.**

794

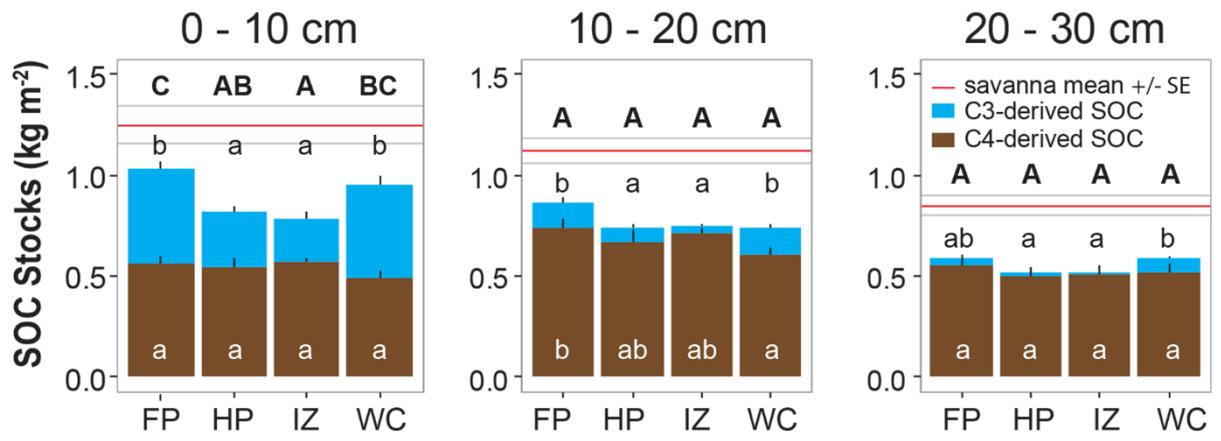


795

796

797 **Fig. 2.**

798

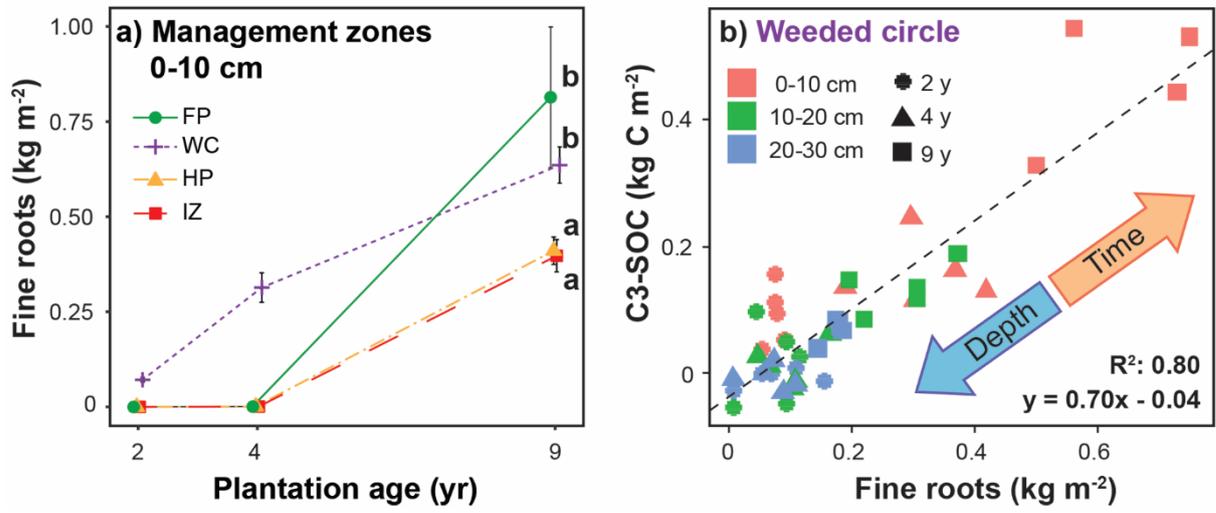


799

800

801 **Fig. 3.**

802

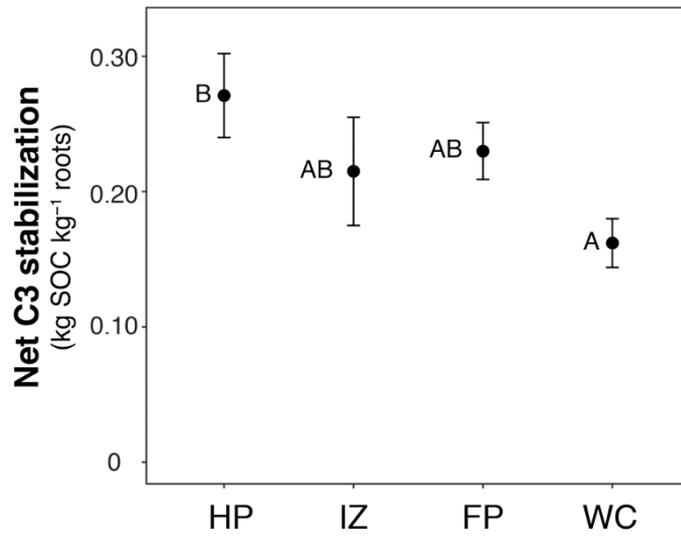


803

804

805 **Fig. 4.**

806

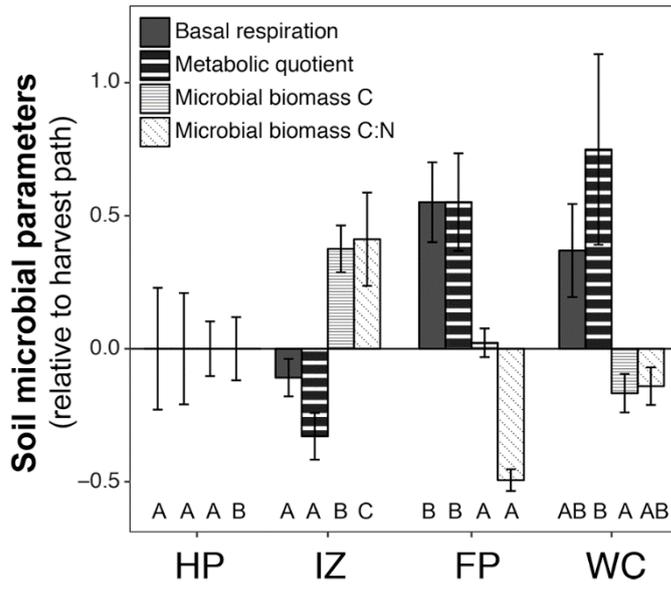


807

808

809 **Fig. 5.**

810

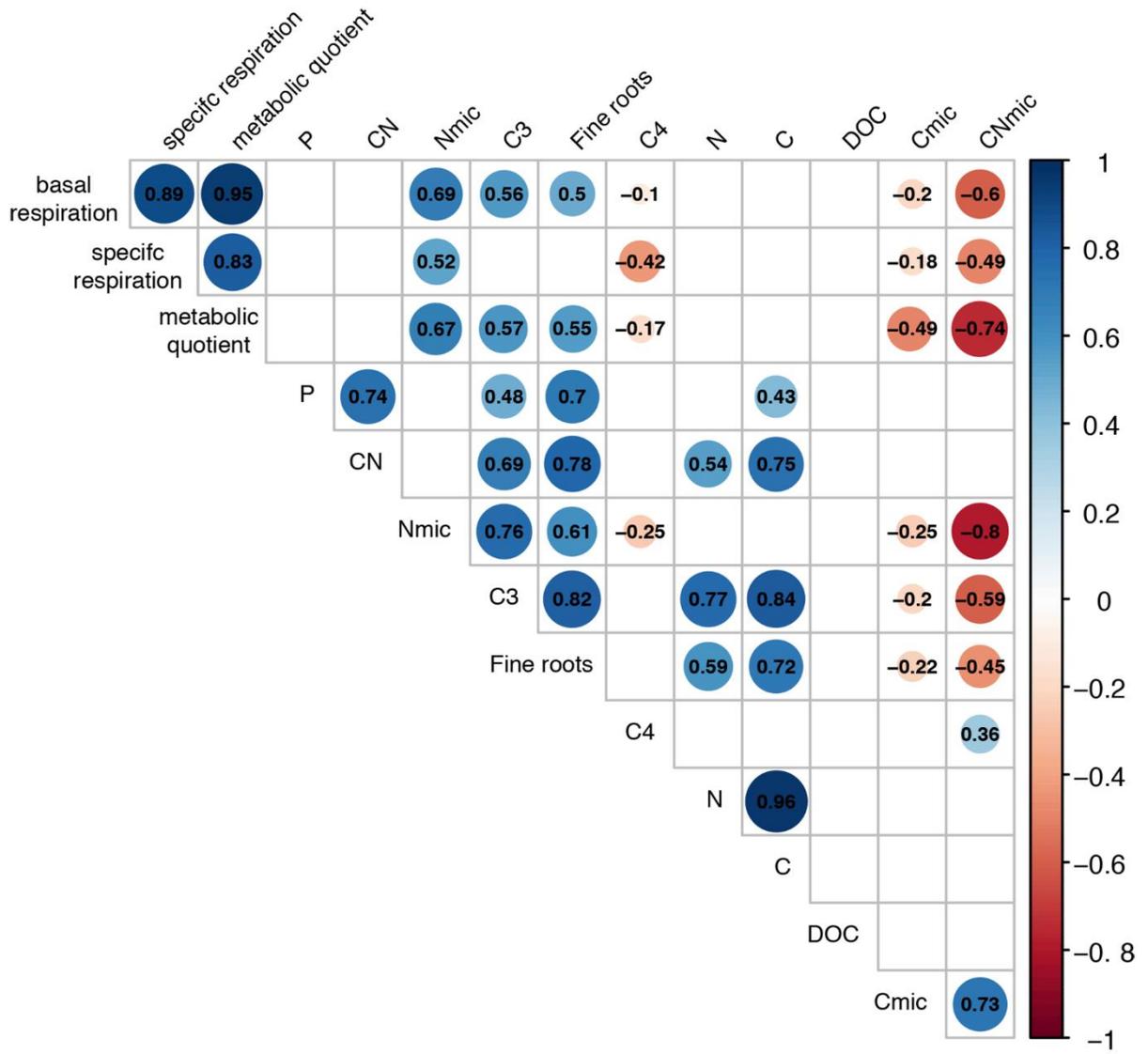


811

812

813 **Fig. 6.**

814

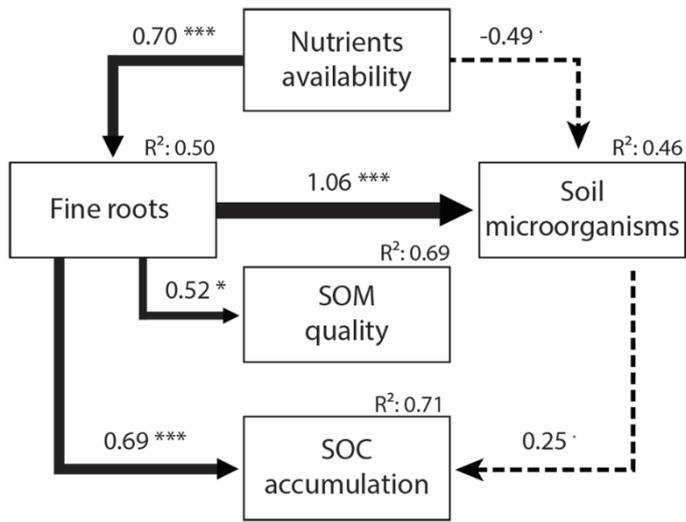


815

816

817 **Fig. 7.**

818



819

820